

## Fortran-77 Computer Programs for 2D and 2.5D analysis of gravity anomalies with variable density contrast

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**Abstract.** Three computer programs in Fortran 77 are presented to calculate the gravity anomalies of 2.5-D and 2-D bodies with either constant density contrast or linear decrease of density contrast with depth or a second order decrease of density contrast with depth. The decrease of density contrast in sedimentary basins may be approximated by a quadratic function. A sedimentary basin may be viewed as a number of 2-D or 2.5-D prisms placed in juxtaposition. Approximate equations for both cases are developed for rapid calculation of gravity anomalies. Efficient methods are developed for two-and-half and two-dimensional analysis of gravity anomalies by an appropriate use of the exact and approximate equations. The depths to the basement are adjusted iteratively by comparing the calculated anomalies with the observed anomalies. A computer program is also developed to deduce the structure of a sedimentary basin along any off-set profile by a suitable modification of the anomaly equations. By interpreting several off-set profiles covering the entire basin, a depth contour map can be drawn which is comparable to a structural map deduced by 3-D methods. Two field examples are given demonstrating the utility of the computer programs.

### 1. Introduction

The interpretation of gravity anomalies may be carried out with the assumption that the causative body is two-dimensional. Many methods are available in the literature to calculate gravity anomalies of two-dimensional bodies. Talwani et al. (1959) developed methods to calculate the gravity anomalies of 2-D bodies of arbitrary shape. Bott (1960) has developed a method for modelling of gravity anomalies over sedimentary basins. But in nature two-dimensional bodies rarely occur and most bodies have finite strike length. There are several algorithms available in the literature for interpretation of gravity anomalies of three-dimensional bodies. Nagy (1966), Talwani and Ewing (1960), Cordell and Henderson (1968) have developed methods for 3-D

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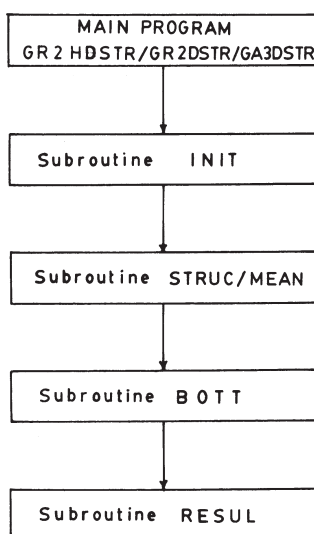


Fig. 1 - The linking of the subroutines with the main program.

modelling of gravity anomalies. Goetze and Lahmeyer (1988) have developed the 3-D interactive modelling for gravity and magnetic anomalies. Many causative bodies can be assumed to have a uniform vertical cross-section and terminate at a finite distance, and these may be called 2-D bodies of finite strike length or 2.5-D bodies. Though the 3-D programs can be used for modelling of gravity anomalies over 2.5-D bodies, they require clearly a lot of computer time. Instead, methods can be developed for modelling the gravity anomalies over such bodies along a profile taken perpendicular to its strike length by deriving suitable equations.

The equations for the gravity anomaly of 2.5-D polygon bodies with constant density contrast were determined by Rasmussen and Pedersen (1979) and a computer program for calculation of these anomalies was presented by Enmark (1981). Bhaskara Rao (1986) has shown that the decrease of density contrast in sedimentary basins can be approximated by a quadratic density function. Bhaskara Rao, Prakash and Ramesh Babu (1990) presented 3-D and 2.5-D modelling techniques of gravity anomalies with a quadratic density function. The approximation of the decrease of density contrast in sedimentary basins by a quadratic function makes it possible to derive a closed form solution in the space domain. The principle is used here to develop the generalized computer programs GR2HDSTR for 2.5-D, and GR2DSTR for 2-D inversion of gravity data either for a constant density contrast or for a variable density contrast with depth. A computer program GA3DSTR is also presented to deduce the structure of a sedimentary basin along any off-set profile, and hence to construct a contour map for basement depths.

## 2. Anomaly equations

The decrease of density contrast in many sedimentary basins could be fairly accurately approximated by a quadratic function (Bhaskara Rao, 1986) as

$$\Delta\rho(z) = a_0 + a_1 z + a_2 z^2, \quad (1)$$

where  $z$  represents the depth measured positive in the downward direction,  $a_0$  represent the extrapolated value of density contrast at the surface, and  $a_1$  and  $a_2$  are the constants of the quadratic function.

For 2.5-D analysis of the gravity anomalies, a sedimentary basin may be viewed as a number of vertical prisms of half thickness  $T$ , and half strike length  $Y$ , symmetrically placed below each observation point along a profile. Thus, the equation for the gravity anomaly of a prismatic model of half strike length  $Y$  is given by Bhaskara Rao et al. (1990) as,

$$\begin{aligned} \Delta g(x) = & 2\gamma a_0 \left[ \left[ z \arctan \frac{YF}{zr} + \frac{F}{2} \ln \frac{r-Y}{r+Y} + \frac{Y}{2} \ln \frac{r-F}{r+F} \right]_{z=Z_1}^{Z_2} \right]_{u=-T}^T \\ & + \frac{2\gamma a_1}{2} \left[ \left[ z^2 \arctan \frac{YF}{zr} + 2YF \ln 2(r+z) - Y^2 \arctan \frac{Fz}{Yr} - F^2 \arctan \frac{Yz}{Fr} \right]_{z=Z_1}^{Z_2} \right]_{u=-T}^T \\ & + \frac{2\gamma a_2}{3} \left[ \left[ z^3 \arctan \frac{YF}{zr} + 2YFr - \frac{F^3}{2} \ln \frac{r-Y}{r+Y} - \frac{Y^3}{2} \ln \frac{r-F}{r+F} \right]_{z=Z_1}^{Z_2} \right]_{u=-T}^T \end{aligned} \quad (2)$$

where  $F = u-x$ ,  $r = \sqrt{(u-x)^2 + z^2 + Y^2}$ , and  $\gamma$  is the gravitational constant. The other parameters are shown in Fig. 2. The corresponding equation for a 2-D prismatic model is given by Bhaskara Rao (1986).

### 3. Approximate equation

As computation of the gravity anomaly at any point involves repeated use of Eq. (2), it necessarily takes a lot of computer time. In order to save computer time and make the programs of practical use, we have developed an approximate equation for anomaly calculation. This is valid for a small distance from the centre of the prism, and is obtained by treating the prism as a sheet mass. Thus, the approximate equation is given by Bhaskara Rao et al. (1990) as,

$$\begin{aligned} \Delta g(x) = & \gamma a_0 \Delta x \left[ \ln \frac{r-Y}{r+Y} \right]_{z=Z_1}^{Z_2} \\ & + 2\gamma a_1 \Delta x \left[ Y \ln 2(r+z) - x \arctan \frac{Yz}{xr} \right]_{z=Z_1}^{Z_2} \\ & + \gamma a_2 \Delta x \left[ 2Yr - x^2 \ln \frac{r-Y}{r+Y} \right]_{z=Z_1}^{Z_2}, \end{aligned} \quad (3)$$

where  $r = \sqrt{x^2 + Y^2 + z^2}$  and  $\Delta x$  is the station spacing. The corresponding equation in 2-D case

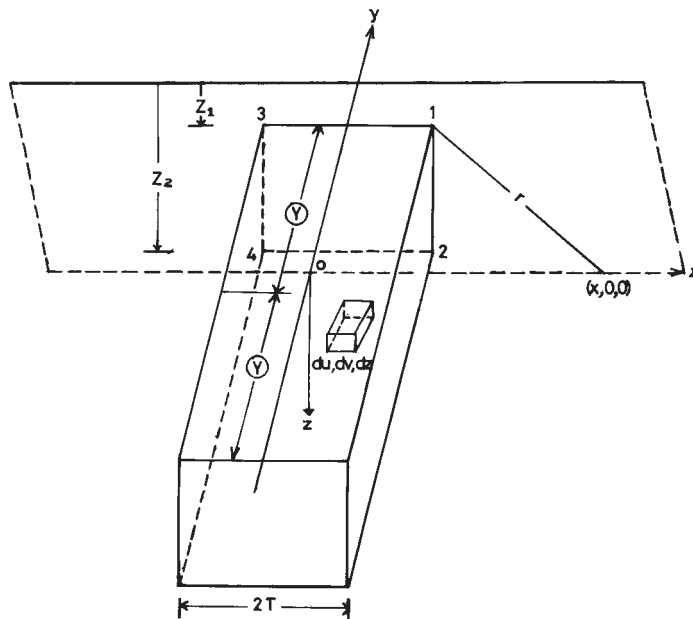


Fig. 2 - 2.5-D Prismatic model.

is given by

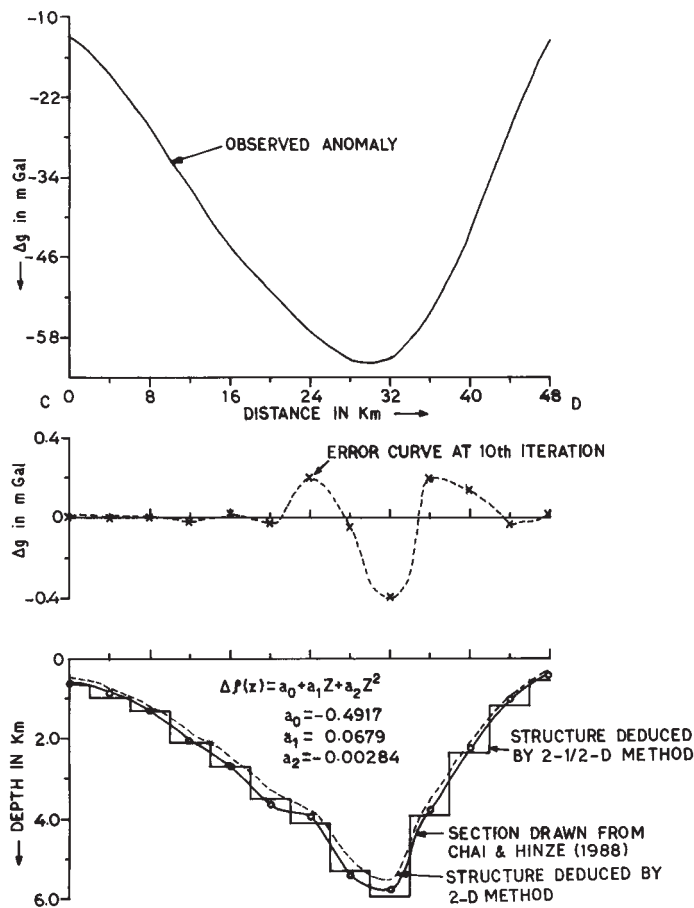
$$\begin{aligned} \Delta g(x) = & \gamma a_0 \Delta x \left[ \ln r \right]_{z=Z_1}^{Z_2} \\ & + 2 \gamma a_1 \Delta x \left[ z - x \arctan \frac{z}{x} \right]_{z=Z_1}^{Z_2} \\ & + \gamma a_2 \Delta x \left[ z^2 - x^2 \ln r \right]_{z=Z_1}^{Z_2}, \end{aligned} \tag{4}$$

where  $r = x^2 + z^2$  and  $\Delta x$  is the station spacing.

### 3. Application to 3-D modelling

Eqs. (2) and (3) are used to deduce the structure of a basin along a profile usually taken perpendicular to the elongation of the anomaly contours and passing through the minimum anomaly point, which may be called the principal profile. With a suitable modification of Eqs. (2) and (3), we can also deduce the structure of a basin along any other profile taken parallel to the principal profile. Thus, the equation for the gravity anomaly of a prismatic model, which is at an off-set distance of  $y$  from the principal profile, may be given by

$$\Delta g(x, y) = \frac{\Delta g(x, Y + y) + \Delta g(x, Y - y)}{2}. \tag{5}$$



**Fig. 3** - Structure derived by 2.5-D modelling of an off-set profile CD (Fig. 4) across the Los Angeles basin using a quadratic density function.

In other words, the gravity anomaly at any point is obtained by taking the average of the gravity anomalies calculated by replacing  $Y$  by  $Y+y$  and  $Y-y$ , respectively, in Eqs. (2) and (3). The gravity anomaly along several profiles taken parallel to the principal profile at uniform off-set intervals covering the entire basin are interpreted by the above method, and the depths of the prisms along these profiles are deduced. The depths are used to construct a depth contour map, which may look like the structural map of a sedimentary basin. This method is comparable to the 3-D methods for deducing the structural map of a basin both in accuracy and computer time.

#### 4. Computer programs

The computer program GR2HDSTR, coded in Fortran 77 for interpretation of gravity anomalies along the principal profile, is given in Appendix 1. This is used to calculate the basement topography of a sedimentary basin using a quadratic density function to account for variation of

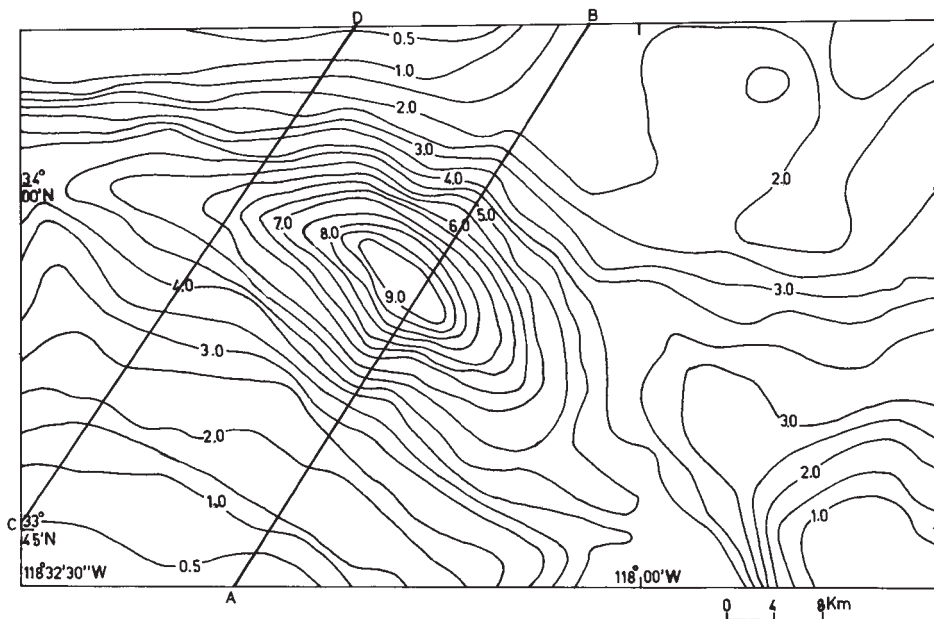


Fig. 4 - Basement contour map of the Los Angeles basin derived from 2.5-D modelling of several off-set gravity profiles using a quadratic density function. Contour interval is 0.5 km.

density contrast with depth. A prism of finite strike length is assumed to be symmetrically placed below each observation point. The data is equispaced, and the width of the prisms is the same as the station interval. The average half strike length,  $Y$ , for each prism may be obtained by measuring the lengths of the elongation of the contours perpendicular to either side of the principal profile.  $Y$  may be different from one prism to another. The input consists of gravity anomalies, station spacing, number of observations, half strike length of each prism, and coefficients of quadratic density function. Total number of iterations and LT, the regions specifying the use of the exact and approximate equations, also may be specified as input.

Fig. 1 shows the linking of the subroutines, and the working of the algorithm is given below.

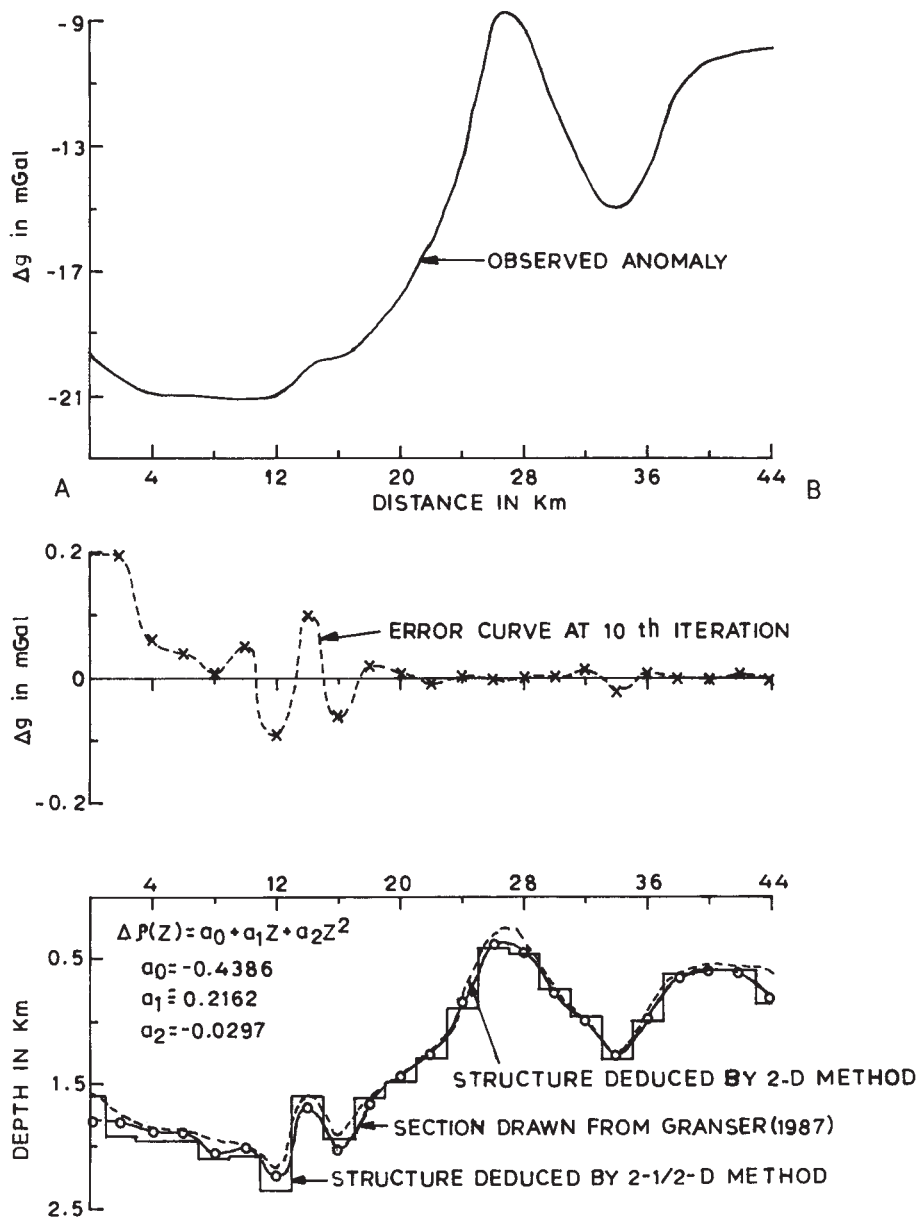
**Main Program:** The main program consists of input and output parameters and calling of various subroutines.

**INIT:** The subroutine subprogram 'INIT' calculates the initial depth to the bottom of the prism by using the Bouguer slab formula  $Z(k) = \Delta g_{obs}(k) / 2\pi\gamma a_0$ .

**STRUC:** The subroutine 'STRUC' calculates the gravity anomalies due to a given model of the sedimentary basin by appropriately using exact and approximate equations ANOM1 and ANOM2, respectively.

**BOTT:** The subroutine subprogram 'BOTT' iteratively improves the depths of the prisms by a method based on the differences between the observed and calculated anomalies. The increment or decrement to be given to the depth is

$$\Delta Z(k) = [\Delta g_{obs}(k) - \Delta g_{cal}(k)] / 2\pi\gamma\Delta\rho(z). \quad (6)$$

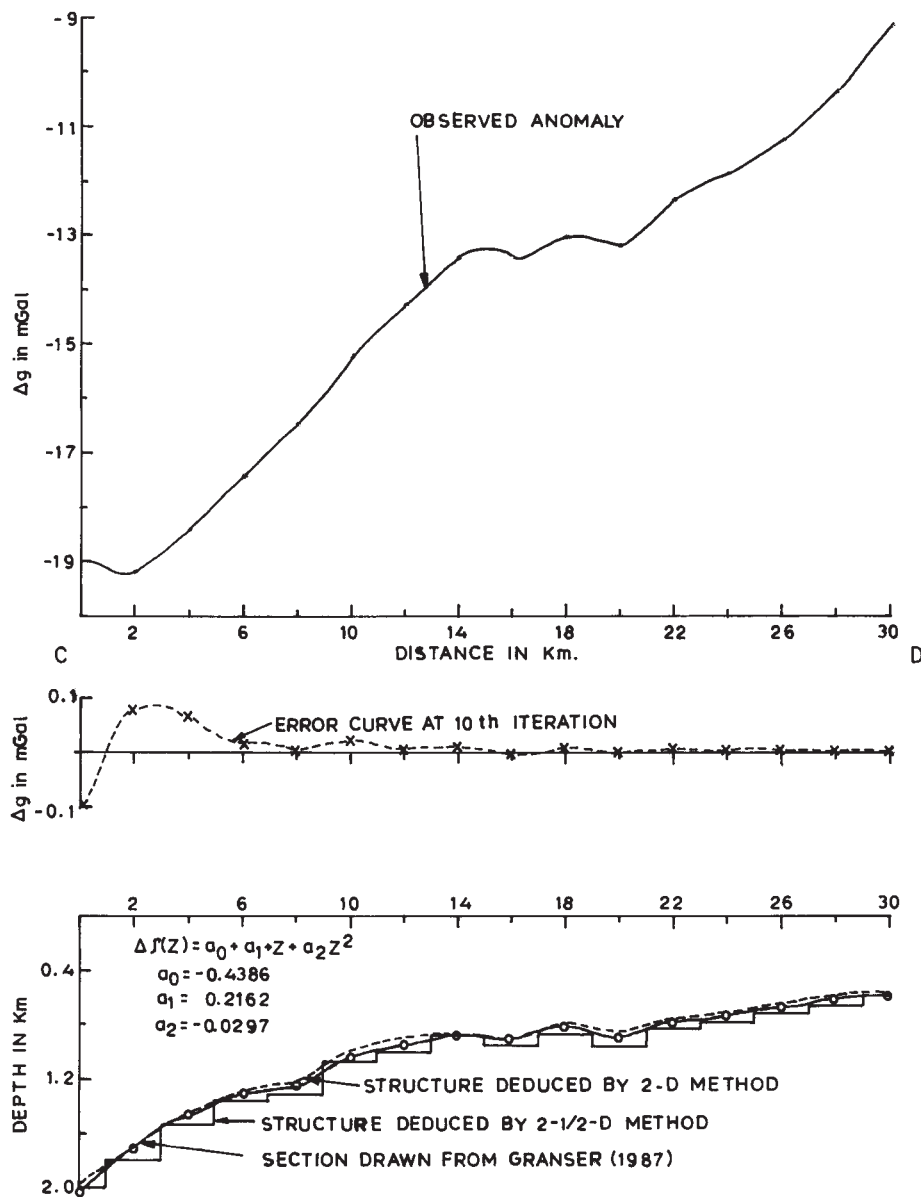


**Fig. 5** - Structure derived by 2.5-D modelling of the principal gravity profile AB (Fig. 7) across the Pannonian basin using a quadratic density function. The corresponding interpretation with 2-D modelling is given by the dotted curve.

ANOM1: The function subprogram 'ANOM1' is used to calculate the gravity anomalies using the exact equation (Eq. (2)).

ANOM2: The function subprogram 'ANOM2' is used to calculate the gravity anomalies using the approximate equation (Eq. (3)).

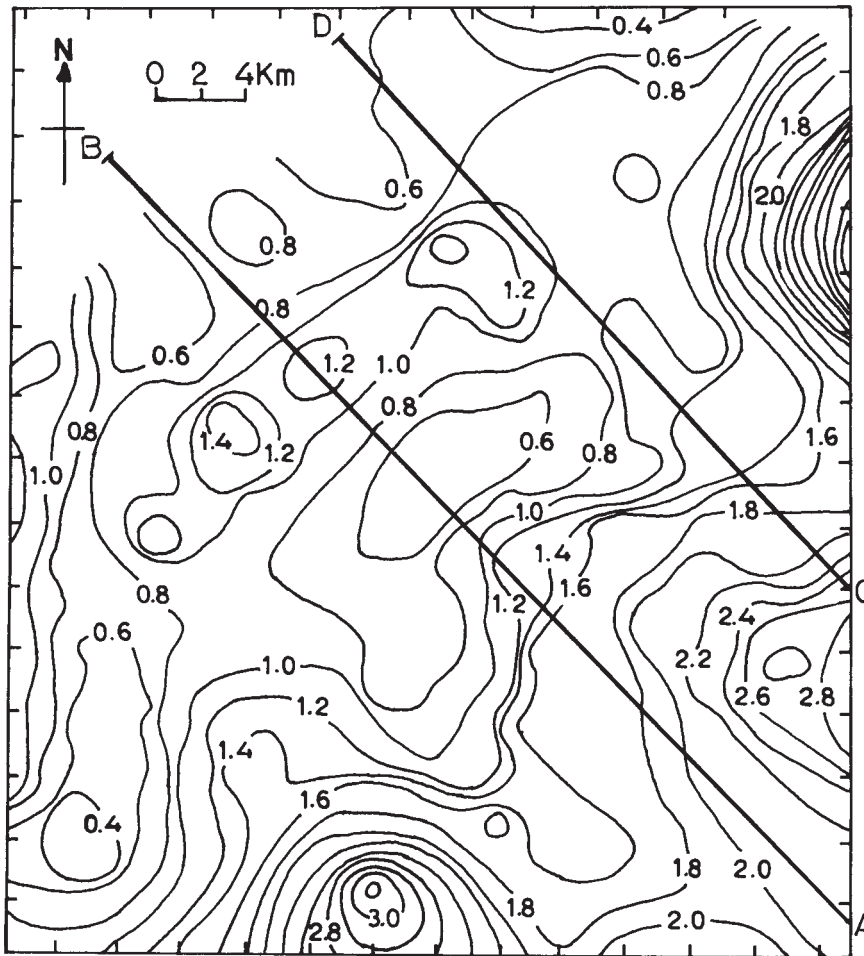
RESUL: The subroutine 'RESUL' gives the results such as the distances from one end of the profile and the corresponding depths to the basement.



**Fig. 6** - Structure derived by 2.5-D modelling of an off-set gravity profile (CD in Fig. 7) across the Pannonian basin using a quadratic density function.

The use of ANOM1 and ANOM2 in each stage is defined by the value of LT we assign. It has the value of LT1 (1) in the first stage, LT1 (2) in the second stage and LT1 (3) in the third stage. ANOM1 may be used to calculate the anomalies on the top and immediate neighbouring points of the prisms [LT1 (1)]. In the intermediate stage, the domain for the use of the exact equation may be extended to three points. In the last stage, the use of the exact equation may be extended to five stations from the centre of the prisms.





**Fig. 7** - Basement contour map derived from 2.5-D modelling of several off-set profiles across the Pannonian basin using a quadratic density function. Contour interval is 0.2 km.

The programs work well for the situation of outcropping prisms by imposing a simple constraint  $Z_1=0.0001$  km, to avoid any possible singularities. These programs can also be used without any modification for the situation of a constant density contrast or for a linear decrease of density contrast with depth by properly assigning zeros to  $a_1$  and  $a_2$ .

The corresponding computer program GR2DSTR for 2-D modelling of sedimentary basins is given in Appendix 2. This program is supported by the subroutines INIT, STRUC, BOTT, ANOM1, ANOM2 and RESULT. The subroutines INIT, STRUC, BOTT and RESULT are the same as given in the program GR2HDSTR, and the listing of the subroutines ANOM1 and ANOM2 are given in Appendix 2.

A computer program GA3DSTR is also developed in Fortran 77 to deduce the structure of a sedimentary basin along any off-set profile and is given in Appendix 3. The only difference in this program from GR2HDSTR is that the values of half strike lengths are modified depending

upon the value of the off-set distances,  $y$ , from the principal profile. Two half strike lengths are assigned to each prism by measuring the lengths of elongation of the contours on either side of the profile. These two values correspond to  $Y+y$  and  $Y-y$  in Eq. (5). The observed anomalies,  $\Delta g_{obs}$ , are measured along the off-set profile, and the calculated anomalies,  $\Delta g_{cal}$ , are obtained using Eq. (2) or Eq. (3), and the structure of the basin along the off-set profile is obtained using the differences between the observed and calculated anomalies as described above. By interpreting several off-set profiles covering the entire basin, we can deduce the 3D structural map of a basin. The program GA3DSTR is supported by a subroutine MEAN, along with the subroutines INIT, STRUC, BOTT and RESUL, and function subprograms ANOM1 and ANOM2 described in the program GR2HDSTR. The subroutine MEAN calculates the average gravity anomalies along any off-set profile using Eq. (5).

## 5. Examples of interpretation

The programs developed in this paper are applied to the interpretation of the residual gravity anomaly contour maps of the Los Angeles basin, California, (Chai and Hinze, 1988) and the Western part of the Pannonian basin, Eastern Austria (Granser, 1987). The density vs. depth data corresponding to the function  $\Delta\rho(z) = -0.5 \exp(-0.1609z)$  given by Chai and Hinze (1988) is fitted to a quadratic function by the least-squares method, and the coefficients are estimated ( $a_0 = -0.49$ ,  $a_1 = 0.067$  and  $a_2 = -0.0028$ ).

The residual gravity contour map published by Chai and Hinze (1988) has been interpreted by the 2.5-D method described in the text by taking the principal profile AB (Fig. 4) and several off-set profiles, such as CD (Fig. 4). The principal profile AB is sampled at equispaced intervals of 4 km, which is the width of each prism. The average half strike length of each prism is measured from the gravity contour map perpendicular to AB. The results of the inversion, and the corresponding error curve at the end of the 10th iteration are given by Bhaskara Rao et al. (1990). The depth section obtained by this method closely fits the section derived from the structural contour map published by Chai and Hinze (1988). The CPU time for 10 iterations on a PC 80386 computer using the program GR2HDSTR is about 1 min. However, if we carry out the inversion exclusively with the exact equation, the time would be about 2 min.

Fig. 3 shows the interpretation of a gravity anomaly profile taken along CD, which is parallel to the principal profile AB, and is at an off-set distance of 16 km, as shown in Fig. 4. This profile is interpreted using GA3DSTR and also by GR2DSTR. The depth section obtained along this profile by the present method and that taken from the structural map given by Chai and Hinze (1988) are also given in Fig. 3. These two interpretations are in close agreement.

The computer program GA3DSTR is also used to deduce the depths to the basement from the gravity anomalies along various off-set profiles taken at a uniform interval of 4 km, and covering the entire basin. Based on these interpreted depths, a depth contour map is drawn and is given in Fig. 4. This contour map is similar to the structural map derived by Chai and Hinze (1988). The advantage of analysing the gravity profiles with the present method is that different strike lengths of the prisms can be taken into account.

Fig. 5 represents another example of the inversion results for the principal profile AB (Fig. 7) in the Western part of the Pannonian basin, Eastern Austria (Granser, 1987). The density vs. depth data corresponding to the function  $\Delta\rho(z) = -0.45 \exp(-0.65z)$  given by Granser (1987) is fitted to a quadratic function by the least-squares method, and the coefficients are estimated ( $a_0 = -0.4386$ ,  $a_1 = 0.2162$  and  $a_2 = -0.0297$ ). The computer program GR2HDSTR is used to deduce the structure of the basin along this profile. The results obtained using the 2-D program GR2DSTR are also shown in Fig. 5, as are the depths to the basement taken from the structural map given by Granser (1987) along this profile. It may be observed that the depths obtained by the present methods agree with those given by Granser (1987).

Fig. 6 shows the interpretation of an off-set gravity anomaly profile taken along profile CD using the computer program GA3DSTR. The profile is at an off-set distance of 12 km from the principal profile as shown in Fig. 7. The anomalies are sampled at an interval of 2 km. Two half strike lengths are measured from the contour map on either side of the profile CD. The interpreted depths of the prisms are shown in Fig. 6. The depths obtained by Granser (1987) along this profile are also drawn by circles with continuous lines for comparison. These two results are in good agreement.

The computer program GA3DSTR is similarly used to deduce the depths to the basement from gravity anomalies along several profiles taken parallel to the principal profile at constant interval of 2 km, and covering the entire basin. The residuals are below  $\pm 0.1$  mGal along any profile. The interpreted depths to the basement are drawn in the form of a depth contour map which is shown in Fig. 7. This map is similar to the structural map given by Granser (1987).

**Acknowledgments.** One of the authors (M. J. Prakash) is grateful to the Council of Scientific and Industrial Research, New Delhi, for providing financial assistance by way of Research Associate.

#### Appendix 1 - Computer program 6R2HDSTR coded in Fortran 77

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*****
*****
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```
PROGRAM GR2HDSTR TO DEDUCE THE 2.5-D STRUCTURE OF A
SEDIMENTARY BASIN FROM ITS GRAVITY ANOMALIES WITH A
QUADRATIC DENSITY FUNCTION.
```

```
INPUT
- - -
```

```
GOBS      : OBSERVED GRAVITY ANOMALIES (MGALS)
NX        : NUMBER OF OBSERVATIONS IN X-DIRECTION
DX        : STATION INTERVAL IN X-DIRECTION (KM)
A0,A1&A2 : COEFFICIENTS OF QUADRATIC DENSITY
           FUNCTION (GM/CC)
ITER1     : TOTAL NUMBER OF ITERATIONS REQUIRED
ITR1      : NUMBER OF ITERATIONS REQUIRED IN FIRST STAGE
ITR2      : NUMBER OF ITERATIONS REQUIRED IN SECOND STAGE
LT1       : LIMITING VALUE OF THE DOMAIN IN WHICH EXACT
           EQUATION IS USED FOR ANOMALY CALCULATION
```

```

      Y      :  HALF STRIKE LENGTHS OF THE PRISMS (KM)

OUTPUT
-_-_-

      GCAL   :  CALCULATED GRAVITY ANOMALIES (MGALS)
      Z      :  DEPTH TO THE BOTTOM (KM)

SUPPORTING SUBROUTINES AND SUBPROGRAMS:
-_-_-_-_-
              INIT, STRUC, BOTT, RESUL, ANOM1, ANOM2
*****
*****
      DIMENSION X(40),GOBS(40),GCAL(40),Z(40),Y(40),LT1(3)
      PI=3.14159265
      CONST=PI*40.0/3.0
      READ(*,801)NX,ITER1,DX,A0,A1,A2
      READ(*,802)ITR1,ITR2
      READ(*,803)(LT1(I),I=1,3)
      WRITE(*,901)NX,ITER1,DX,A0,A1,A2
      WRITE(*,902)ITR1,ITR2
      WRITE(*,903)(LT1(I),I=1,3)
      DO 10 I=1,NX
10      X(I)=FLOAT(I-1)*DX
      READ(*,804)(GOBS(I),I=1,NX)
      WRITE(*,904)(GOBS(I),I=1,NX)
      READ(*,804)(Y(I),I=1,NX)
      CALL INIT(NX,A0,X,GOBS,CONST,Z)
      WRITE(*,905)(Z(I),I=1,NX)
      WRITE(*,905)(Y(I),I=1,NX)
      DO 20 ITER = 1,ITER1
          CALL STRUC(NX,DX,A0,A1,A2,X,Y,Z,ITER,ITR1,ITR2,LT1,GCAL)
          CALL RESUL(NX,ITER,X,Z,GOBS,GCAL)
          CALL BOTT(NX,CONST,A0,A1,A2,GOBS,GCAL,Z)
20      CONTINUE
801      FORMAT(2I5,4F10.4)
802      FORMAT(2I5)
803      FORMAT(3I5)
804      FORMAT(8F10.4)
901      FORMAT(2I5,4F10.4)
902      FORMAT(2I5)
903      FORMAT(3I5)
904      FORMAT(8F10.4)
905      FORMAT(/8F10.4)
      STOP
      END
*****
*****
      SUBROUTINE INIT :  CALCULATES THE INITIAL ESTIMATES OF
                        THE STRUCTURE.

INPUT
-_-_-
      NX      :  NUMBER OF OBSERVATIONS IN X-DIRECTION
      A0      :  CONSTANT DENSITY CONTRAST (GM/CC)
      X      :  DISTANCE TO THE ANOMALY POINT IN X-DIRECTION

```

```

          GOBS      :   OBSERVED GRAVITY ANOMALIES (MGALS)
          CONST     :   PI*40.0/3.0

      OUTPUT
      - - - -
          Z          :   DEPTH TO THE BOTTOM (KM)
*****
*****
      SUBROUTINE INIT(NX,A0,X,GOBS,CONST,Z)
      DIMENSION X(40),GOBS(40),Z(40)
      DO 10 I=1,NX
          Z(I)=GOBS(I)/(CONST*A0)
          IF (Z(I).LE.0.0) THEN
              Z(I)=0.00001
          ENDIF
10      CONTINUE
      RETURN
      END
*****
*****
      SUBROUTINE STRUC: CALCULATES THE GRAVITY ANOMALIES OF A
                      SEDIMENTARY BASIN BY THE COMBINED
                      USE OF EXACT AND APPROXIMATE EQUATIONS
                      OF PRISMATIC MODEL.

      INPUT
      - - - -
          NX          :   NUMBER OF OBSERVATIONS IN X-DIRECTION
          DX          :   STATION INTERVAL IN X-DIRECTION (KM)
          A0,A1      :   COEFFICIENTS OF THE QUADRATIC DENSITY
          & A2      :   FUNCTION (GM/CC)
          X          :   DISTANCE TO THE ANOMALY POINT IN X-
DIRECTION
          Z          :   DEPTH TO THE BOTTOM (KM)
          ITER       :   ITERATION NUMBER
          ITR1       :   NUMBER OF ITERATIONS REQUIRED IN FIRST
STAGE
          ITR2       :   NUMBER OF ITERATIONS REQUIRED IN SECOND
STAGE
          LT1        :   LIMITING VALUE OF THE DOMAIN IN WHICH
EXACT
                      EQUATION IS USED FOR ANOMALY CALCULATION
          Y          :   HALF STRIKE LENGTH OF THE PRISM (KM)

      OUTPUT
      - - - -
          GCAL       :   CALCULATED GRAVITY ANOMALIES (MGALS)
*****
*****
      SUBROUTINE STRUC(NX,DX,A0,A1,A2,X,Y,Z,ITER,ITR1,ITR2,LT1,GCAL)
      DIMENSION X(40),Z(40),GCAL(40),Y(40),LT1(3)
      DO 10 I=1,NX
10      GCAL(I)=0.0
          Z1=0.0001
          IF (ITER.LT.ITR1) THEN
              LT=LT1(1)

```

```

      GOTO 50
    ENDIF
    IF (ITER.LE.ITR2) THEN
      LT=LT1(2)
      GOTO 50
    ELSE
      LT=LT1(3)
    ENDIF
50   DO 20 I=1,NX
      DO 20 K=1,NX
        X1=X(I)-X(K)
        Z2=Z(K)
        Y1=Y(K)
        IF (ABS(I-K).LE.LT) THEN
          GCAL1=ANOM1(DX,Y1,X1,Z1,Z2,A0,A1,A2)
          GOTO 100
        ELSE
          GCAL1=ANOM2(DX,Y1,X1,Z1,Z2,A0,A1,A2)
        ENDIF
100   GCAL(I)=GCAL(I)+GCAL1
20   CONTINUE
      RETURN
    END
*****
      SUBROUTINE BOTT: CALCULATES THE STRUCTURE OF A
                    BASIN USING THE
                    BOTT'S METHOD.

      INPUT
      - - -
          NX          :      NUMBER OF OBSERVATIONS IN X-DIRECTION
          A0,A1       :      COEFFICIENTS OF QUADRATIC DENSITY
          & A2        :      FUNCTION (GM/CC)
          GOBS        :      OBSERVED GRAVITY ANOMALIES (MGALS)
          GCAL        :      CALCULATED GRAVITY ANOMALIES (MGALS)
          CONST       :      PI*40.0/3.0

      OUTPUT
      - - -
          Z           :      DEPTH TO THE BASEMENT (KM)
*****
      SUBROUTINE BOTT(NX,CONST,A0,A1,A2,GOBS,GCAL,Z)
      DIMENSION GOBS(40),GCAL(40),Z(40),DZ(40)
      DO 10 I=1,NX
        R2=A0+A1*Z(I)+A2*Z(I)*Z(I)
        CONST1=CONST*R2
10   DZ(I)=(GOBS(I)-GCAL(I))/CONST1
      DO 20 I=1,NX
        Z(I)=Z(I)+DZ(I)
        IF (Z(I).LE.0.0) THEN
          Z(I)=0.00001
        ENDIF
20   CONTINUE
      RETURN

```

```

END
*****
*****
FUNCTION SUBPROGRAM ANOM1: CALCULATES THE ANOMALY OF
PRISMATIC MODEL WITH EXACT
EQUATION.

INPUT
- - -
          DX      : STATION INTERVAL IN X-DIRECTION (KM)
          A0,A1   : COEFFICIENTS OF QUADRATIC DENSITY
                   & A2             FUNCTION (GM/CC)
          X1      : DISTANCE TO THE POINT OF CALCULATION
                   IN X-DIRECTION WITH REFERENCE TO THE
RATION      EPICENTRE OF THE PRISM UNDER CONSIDERATION

          Y1      : HALF STRIKE LENGTH OF THE PRISM (KM)
          Z1      : DEPTH TO THE TOP OF THE PRISM (KM)
          Z2      : DEPTH TO THE BOTTOM OF THE PRISM (KM)
          T       : HALF THICKNESS OF THE PRISM (KM)
          GAMA    : GRAVITATIONAL CONSTANT
*****
*****
FUNCTION ANOM1 (DX, Y1, X1, Z1, Z2, A0, A1, A2)
T=DX/2.0
GAMA=20.0/3.0
F1=X1-T
F2=X1+T
IF (F1.EQ.0.0) THEN
  F1=0.00001
ENDIF
IF (F2.EQ.0.0) THEN
  F2=0.00001
ENDIF
IF (Z1.EQ.0.0) THEN
  Z1=0.00001
ENDIF
R1=SQRT (Z1*Z1+F1*F1+Y1*Y1)
R2=SQRT (Z2*Z2+F1*F1+Y1*Y1)
R3=SQRT (Z1*Z1+F2*F2+Y1*Y1)
R4=SQRT (Z2*Z2+F2*F2+Y1*Y1)
T1=ATAN ( (Y1*F1) / (Z1*R1) ) - ATAN ( (Y1*F2) / (Z1*R3) )
T2=ATAN ( (Y1*F2) / (Z2*R4) ) - ATAN ( (Y1*F1) / (Z2*R2) )
T3=ALOG ( ( (R2+F1) * (R3+F2) ) / ( (R1+F1) * (R4+F2) ) )
T4=ALOG ( ( (R2+Y1) * SQRT (Z1*Z1+F1*F1) ) / ( (R1+Y1) * SQRT (Z2*Z2+F1*F1) ) )
T5=ALOG ( ( (R3+Y1) * SQRT (Z2*Z2+F2*F2) ) / ( (R4+Y1) * SQRT (Z1*Z1+F2*F2) ) )
T6=ALOG ( (R1+Z1) / (R2+Z2) )
T7=ATAN ( (F1*Z2) / (Y1*R2) ) - ATAN ( (F1*Z1) / (Y1*R1) )
T8=ATAN ( (F1*R1) / (Y1*Z1) ) - ATAN ( (F1*R2) / (Y1*Z2) )
T9=ALOG ( (R4+Z2) / (R3+Z1) )
T10=ATAN ( (F2*Z2) / (Y1*R4) ) - ATAN ( (F2*Z1) / (Y1*R3) )
T11=ATAN ( (F2*R3) / (Y1*Z1) ) - ATAN ( (F2*R4) / (Y1*Z2) )
P1=2*GAMA* (Z1*T1+Z2*T2+Y1*T3+F1*T4+F2*T5)
P2=GAMA* ( (Z1*Z1) *T1+ (Z2*Z2) *T2)
P31=2*Y1*F1*T6+Y1*Y1*T7+F1*F1*T8
P3=GAMA*P31

```

```

P41=2*Y1*F2*T9-Y1*Y1*T10-F2*F2*T11
P4=GAMA*P41
P51=((Z1**3)*T1/3.0+(Z2**3)*T2/3.0)
P5=2*GAMA*P51
P61=(2*Y1*(R4-R3)/3.0-(F2*F2)*T5/3.0)
P6=2*GAMA*F2*P61
P71=(2*Y1*(R2-R1)/3.0+(F1*F1)*T4/3.0)
P7=2*GAMA*F1*P71
P8=2*GAMA*Y1**3*(T3/3.0)
TA0=P1
TA1=P2+P3+P4
TA2=P5+P6-P7-P8
ANOM1=A0*TA0+A1*TA1+A2*TA2
RETURN
END

```

```

*****
*****

```

```

FUNCTION SUBPROGRAM ANOM2: CALCULATES THE ANOMALY OF
PRISMATIC MODEL WITH
APPMATE EQUATION.

```

```

INPUT
- - -

```

```

DX      : STATION INTERVAL IN X-DIRECTION (KM)
A0,A1   : COEFFICIENTS OF QUADRATIC DENSITY
& A2    : FUNCTION (GM/CC)
X1      : DISTANCE TO THE POINT OF CALCULATION
          IN X-DIRECTION WITH REFERENCE TO THE
          EPICENTRE OF THE PRISM UNDER CONSIDERATION

```

```

RATION

```

```

Z1      : DEPTH TO THE TOP OF THE PRISM (KM)
Z2      : DEPTH TO THE BOTTOM OF THE PRISM (KM)
Y1      : HALF STRIKE LENGTH OF THE PRISM (KM)
GAMA    : GRAVITATIONAL CONSTANT

```

```

*****
*****

```

```

FUNCTION ANOM2 (DX, Y1, X1, Z1, Z2, A0, A1, A2)
GAMA=20.0/3.0
CONST=2.0*GAMA*Y1*DX
IF (X1.EQ.0.0) THEN
    X1=0.1
ENDIF
R1=X1*X1+Y1*Y1+Z1*Z1
R1=SQRT(R1)
P5=R1-Y1
IF (ABS(P5).LE.0.1) THEN
    P5=0.15
ENDIF
R2=X1*X1+Y1*Y1+Z2*Z2
R2=SQRT(R2)
P1=ALOG((R2-Y1)*(R1+Y1)/((R2+Y1)*P5))
P2=ALOG((R2+Z2)/(R1+Z1))
P3=ATAN(Y1*Z2/(X1*R2))-ATAN(Y1*Z1/(X1*R1))
TP1=CONST*A0*P1/(2.0*Y1)
TP2=CONST*A1*(P2-(X1/Y1)*P3)
TP3=CONST*A2*(R2-R1-X1*X1*P1/(2.0*Y1))

```



```

      ANOM2=TP1+TP2+TP3
      RETURN
    END
*****
*****
      SUBROUTINE RESUL: GIVES THE RESULTS SUCH AS DEPTHS
                      TO THE BASEMENT AND RESIDUAL ANOMALIES.

      INPUT
      _ _ _
          NX          :    NUMBER OF OBSERVATIONS IN X-DIRECTION
          ITER        :    ITERATION NUMBER
          X           :    DISTANCE TO THE ANOMALY POINT IN X-
DIRECTION
          Z           :    DEPTH TO THE BASEMENT (KM)
          GOBS        :    OBSERVED GRAVITY ANOMALIES (MGALS)

      OUTPUT
      _ _ _
          GCAL        :    CALCULATED GRAVITY ANOMALIES (MGALS)
*****
*****
      SUBROUTINE RESUL(NX,ITER,X,Z,GOBS,GCAL)
      DIMENSION X(40),GOBS(40),GCAL(40),RESD(40),Z(40)
      DO 10 I=1,NX
10      RESD(I)=GOBS(I)-GCAL(I)
      WRITE(*,906) ITER
      WRITE(*,907)
      DO 20 I=1,NX
          WRITE(*,908) X(I),Z(I),GOBS(I),GCAL(I),RESD(I)
20      CONTINUE
906      FORMAT(/5X,'ITERATION NO.=' ,I4)
907      FORMAT(/7X,'DISTANCE',9X,'DEPTH',8X,'OBS ANOM',7X,'CAL ANOM'
          *,10X,'RESD',/)
908      FORMAT(5(5X,F10.4))
      RETURN
      END

```

## Appendix 2 - Computer program GR2DSTR coded in Fortran 77

```

*****
*****
      PROGRAM GR2DSTR TO DEDUCE THE 2-D STRUCTURE OF A
      SEDIMENTARY BASIN FROM ITS GRAVITY ANOMALIES WITH
      A QUADRATIC DENSITY FUNCTION.

      INPUT
      _ _ _
          GOBS        :    OBSERVED GRAVITY ANOMALIES (MGALS)
          NX          :    NUMBER OF OBSERVATIONS IN X-DIRECTION
          DX          :    STATION INTERVAL IN X-DIRECTION (KM)
          A0,A1&A2    :    COEFFICIENTS OF QUADRATIC DENSITY

```

```

                FUNCTION (GM/CC)
ITER1          : TOTAL NUMBER OF ITERATIONS REQUIRED
ITR1           : NUMBER OF ITERATIONS REQUIRED IN FIRST STAGE
ITR2           : NUMBER OF ITERATIONS REQUIRED IN SECOND STAGE
LT1           : LIMITING VALUE OF THE DOMAIN IN WHICH EXACT
                EQUATION IS USED FOR ANOMALY CALCULATION

OUTPUT
- - - -

GCAL           : CALCULATED GRAVITY ANOMALIES (MGALS)
Z             : DEPTH TO THE BOTTOM (KM)

SUPPORTING SUBROUTINES AND SUBPROGRAMS:
- - - - -
                INIT, STRUC, BOTT, RESUL, ANOM1, ANOM2
*****
*****
DIMENSION X(40),GOBS(40),GCAL(40),Z(40),LT1(3)
PI=3.14159265
CONST=PI*40.0/3.0
READ(*,801)NX,ITER1,DX,A0,A1,A2
READ(*,802)ITR1,ITR2
READ(*,803)(LT1(I),I=1,3)
WRITE(*,901)NX,ITER1,DX,A0,A1,A2
WRITE(*,902)ITR1,ITR2
WRITE(*,903)(LT1(I),I=1,3)
DO 10 I=1,NX
10  X(I)=FLOAT(I-1)*DX
    READ(*,804)(GOBS(I),I=1,NX)
    WRITE(*,904)(GOBS(I),I=1,NX)
    CALL INIT(NX,A0,X,GOBS,CONST,Z)
    WRITE(*,905)(Z(I),I=1,NX)
    DO 20 ITER = 1,ITER1
        CALL STRUC(ITER,ITR1,ITR2,LT1,NX,DX,A0,A1,A2,X,Z,GCAL)
        CALL RESUL(NX,ITER,X,Z,GOBS,GCAL)
        CALL BOTT(NX,CONST,A0,A1,A2,GOBS,GCAL,Z)
20  CONTINUE
801  FORMAT(2I5,4F10.4)
802  FORMAT(2I5)
803  FORMAT(3I5)
804  FORMAT(8F10.4)
901  FORMAT(2I5,4F10.4)
902  FORMAT(2I5)
903  FORMAT(3I5)
904  FORMAT(8F10.4)
905  FORMAT(8F10.4)
STOP
END
*****
*****
                FUNCTION SUBPROGRAM ANOM1: CALCULATES THE ANOMALYOF
                2-D PRISMATIC MODEL WITH EXACT EQUATION.

INPUT
- - -

```

```

          DX      :      STATION INTERVAL IN X-DIRECTION (KM)
          A0,A1   :      COEFFICIENTS OF QUADRATIC DENSITY
          & A2    :      FUNCTION (GM/CC)
          X1      :      DISTANCE TO THE POINT OF CALCULATION
                       IN X-DIRECTION WITH REFERENCE TO THE
RATION      EPICENTRE OF THE PRISM UNDER CONSIDERATION
          Z1      :      DEPTH TO THE TOP      OF THE PRISM (KM)
          Z2      :      DEPTH TO THE BOTTOM OF THE PRISM (KM)
          T       :      HALF THICKNESS OF THE PRISM (KM)
          GAMA    :      GRAVITATIONAL CONSTANT

```

```

*****
*****

```

```

FUNCTION ANOM1 (DX,X1,Z1,Z2,A0,A1,A2)

```

```

PI=3.14159265

```

```

T=DX/2.0

```

```

GAMA=20.0/3.0

```

```

CONST=2*GAMA

```

```

F1=X1-T

```

```

F2=X1+T

```

```

IF (F1 .EQ. 0.0) THEN

```

```

    F1=0.0001

```

```

ENDIF

```

```

IF (F2 .EQ. 0.0) THEN

```

```

    F2=0.0001

```

```

ENDIF

```

```

IF (Z1 .EQ. 0.0) THEN

```

```

    Z1=0.0001

```

```

ENDIF

```

```

R1=SQRT (Z1*Z1+F2*F2)

```

```

R2=SQRT (Z1*Z1+F1*F1)

```

```

R3=SQRT (Z2*Z2+F1*F1)

```

```

R4=SQRT (Z2*Z2+F2*F2)

```

```

R41=ALOG (R4/R1)

```

```

R32=ALOG (R3/R2)

```

```

PHI1=PI/2+ATAN (F2/Z1)

```

```

PHI2=PI/2+ATAN (F1/Z1)

```

```

PHI3=PI/2+ATAN (F1/Z2)

```

```

PHI4=PI/2+ATAN (F2/Z2)

```

```

PHI43=PHI4-PHI3

```

```

PHI12=PHI1-PHI2

```

```

PHI23=PHI2-PHI3

```

```

PHI14=PHI1-PHI4

```

```

TP1=Z2*PHI43-Z1*PHI12+F2*R41-F1*R32

```

```

TP1=TP1*CONST*A0

```

```

TP2=Z2*Z2*PHI43-Z1*Z1*PHI12+F1*F1*PHI23-F2*F2*PHI14+2*T*(Z2-Z1)

```

```

TP2=TP2*CONST*A1/2.0

```

```

TP3=Z2**3*PHI43-Z1**3*PHI12+F1**3*R32-F2**3*R41+T*(Z2*Z2-Z1*Z1)

```

```

TP3=TP3*CONST*A2/3.0

```

```

ANOM1=TP1+TP2+TP3

```

```

RETURN

```

```

END

```

```

*****
*****

```

```

*****
*****      FUNCTION SUBPROGRAM ANOM2: CALCULATES THE ANOMALY OF
              2-D PRISMATIC MODEL WITH

```

## APPROXIMATE EQUATION.

```

*      _ _ INPUT      *
      DX      :      STATION INTERVAL IN X-DIRECTION (KM)
      A0,A1   :      COEFFICIENTS OF QUADRATIC DENSITY
      & A2    :      FUNCTION (GM/CC)
      X1      :      DISTANCE TO THE POINT OF CALCULATION
                   IN X-DIRECTION WITH REFERENCE TO THE
                   EPICENTRE OF THE PRISM UNDER CONSIDERATION
      Z1      :      DEPTH TO THE TOP      OF THE PRISM (KM)
      Z2      :      DEPTH TO THE BOTTOM OF THE PRISM (KM)
      GAMA    :      GRAVITATIONAL CONSTANT
*****
*****
      FUNCTION ANOM2 (DX,X1,Z1,Z2,A0,A1,A2)
      GAMA=20.0/3.0
      CONST=2.0*GAMA*DX
      IF (X1 .EQ. 0.0) THEN
          X1=0.0001
      ENDIF
      R1=X1*X1+Z1*Z1
      R1=SQRT (R1)
      R2=X1*X1+Z2*Z2
      R2=SQRT (R2)
      R21=ALOG (R2/R1)
      PHI21=ATAN (Z2/X1) -ATAN (Z1/X1)
      TP1=CONST*A0*R21
      TP2=CONST*A1* ((Z2-Z1) -X1*PHI21)
      TP3=CONST*A2* ((Z2*Z2-Z1*Z1) *0.5-X1*X1*R21)
      ANOM2=TP1+TP2+TP3
      RETURN
      END

```

## Appendix 3 - Computer program GA3DSTR coded in Fortran 77

```

*****
*****
      PROGRAM GA3DSTR TO DEDUCE THE 2.5-D STRUCTURE OF A
      SEDIMENTARY BASIN ALONG ANY OFF-SET PROFILE FROM ITS
      GRAVITY ANOMALIES WITH A QUADRATIC DENSITY FUNCTION.

      INPUT
      _ _ _
      GOBS      :      OBSERVED GRAVITY ANOMALIES (MGALS)
      NX        :      NUMBER OF OBSERVATIONS IN X-DIRECTION
      DX        :      STATION INTERVAL IN X-DIRECTION (KM)
      DY        :      SCALE CONSTANT IN Y-DIRECTION (KM)
      A0,A1&A2  :      COEFFICIENTS OF QUADRATIC DENSITY
                   FUNCTION (GM/CC)
      ITER1     :      TOTAL NUMBER OF ITERATIONS REQUIRED

```

```

ITR1      :      NUMBER OF ITERATIONS REQUIRED IN FIRST STAGE
ITR2      :      NUMBER OF ITERATIONS REQUIRED IN SECOND STAGE
LT1       :      LIMITING VALUE OF THE DOMAIN IN WHICH EXACT
              EQUATION IS USED FOR ANOMALY CALCULATION
YY        :      STRIKE LENGTHS ON EITHER SIDE OF THE PROFILE
              CORRESPONDING TO Y+y AND Y-y

```

```

OUTPUT
- - - -

```

```

GCAL      :      CALCULATED GRAVITY ANOMALIES (MGALS)
Z         :      DEPTH TO THE BOTTOM (KM)

```

```

SUPPORTING SUBROUTINES AND SUBPROGRAMS:
- - - - -

```

```

MEAN, INIT, STRUC, BOTT, RESULT, ANOM1, ANOM2

```

```

*****
*****      DIMENSION X(40), GOBS(40), GCAL(40), Z(40), YY(2,40), LT1(3)
*****

```

```

PI=3.14159265
CONST=PI*40.0/3.0
READ(*,801)NX, ITER1,DX,DY,A0,A1,A2
READ(*,802)ITR1, ITR2
READ(*,803)(LT1(I), I=1,3)
WRITE(*,901)NX, ITER1,DX,DY,A0,A1,A2
WRITE(*,902)ITR1, ITR2
WRITE(*,903)(LT1(I), I=1,3)
DO 10 I=1,NX
10  X(I)=FLOAT(I-1)*DX
    READ(*,804)(GOBS(I), I=1,NX)
    WRITE(*,904)(GOBS(I), I=1,NX)
    DO 20 I=1,2
    READ(*,804)(YY(I,J), J=1,NX)
20  CONTINUE
    DO 30 I=1,2
        DO 30 J=1,NX
    YY(I,J)=YY(I,J)*DY
30  CONTINUE
    CALL INIT(NX,A0,X,GOBS,CONST,Z)
    WRITE(*,905)(Z(I), I=1,NX)
    DO 40 I=1,2
    WRITE(*,905)(YY(I,J), J=1,NX)
40  CONTINUE
    DO 50 ITER = 1, ITER1
        CALL MEAN(NX,DX,A0,A1,A2,X,YY,Z,ITER, ITR1, ITR2,LT1,GCAL)
        CALL RESULT(NX,ITER,X,Z,GOBS,GCAL)
        CALL BOTT(NX,CONST,A0,A1,A2,GOBS,GCAL,Z)
50  CONTINUE
801  FORMAT(2I5,5F10.4)
802  FORMAT(2I5)
803  FORMAT(3I5)
804  FORMAT(8F10.4)
901  FORMAT(2I5,5F10.4)
902  FORMAT(2I5)
903  FORMAT(3I5)
904  FORMAT(8F10.4)
905  FORMAT(/8F10.4)

```

```

STOP
END
*****
*****
SUBROUTINE MEAN: CALCULATES THE MEAN GRAVITY ANOMALY
                  USING STRIKE LENGTHS OF Y+y AND Y-y

INPUT
  GOBS          : OBSERVED GRAVITY ANOMALIES (MGALS)
  NX            : NUMBER OF OBSERVATIONS IN X-DIRECTION
  DX            : STATION INTERVAL IN X-DIRECTION (KM)
  A0,A1        : COEFFICIENTS OF THE QUADRATIC DENSITY
  & A2         : FUNCTION (GM/CC)
  ITER          : ITERATION NUMBER
  ITR1          : NUMBER OF ITERATIONS REQUIRED IN FIRST
STAGE
  ITR2          : NUMBER OF ITERATIONS REQUIRED IN SECOND
STAGE
  LT1          : LIMITING VALUE OF THE DOMAIN IN WHICH EXACT
                  EQUATION IS USED FOR ANOMALY CALCULATION
  X            : DISTANCE TO THE ANOMALY POINT IN X-
DIRECTION
  YY           : STRIKE LENGTHS ON EITHER SIDE OF THE
PROFILE
                  CORRESPONDING TO Y+y AND Y-y
  Z            : DEPTH TO THE BOTTOM (KM)

OUTPUT
  GCAL          : CALCULATED GRAVITY ANOMALIES (MGALS)
*****
*****
SUBROUTINE MEAN (NX,DX,A0,A1,A2,X,YY,Z,ITER,ITR1,ITR2,LT1,GCAL)
DIMENSION X(40),Z(40),Y(40),YY(2,40),GCAL(40),GG(2,40),G1(40)
DO 10 I=1,2
  DO 20 J=1,NX
    Y(J)=YY(I,J)
    CALL STRUC (NX,DX,A0,A1,A2,X,Y,Z,ITER,ITR1,ITR2,LT1,G1)
    DO 30 K=1,NX
      GG(I,K)=G1(K)
    30
  10 CONTINUE
  DO 40 I=1,NX
    GCAL (I) =(GG(1,I)+GG(2,I))/2.0
  40 CONTINUE
RETURN
END

```

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